

More on the Gravitational Potential

● Use of Potentials vs Forces / General Comments

- We can think of the potential energy on an object of mass m as $V(\vec{r}) = m\Phi(\vec{r})$, so $\vec{F} = -m\vec{\nabla}\Phi$

• Same idea as electrostatic case

• Then

$$\Phi(\vec{r}) = -\int d^3r' \frac{G\rho(\vec{r}')}{|\vec{r}-\vec{r}'|}$$

is the integral over the mass distribution that acts on m

- You can also add a potential corresponding to the centrifugal force

- Potential determines shape of objects held together by gravity

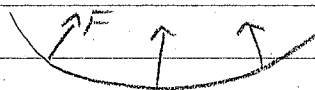
• Consider a fluid surface,

+ the force on fluid is \perp to

surface in equilibrium, or else fluid would flow

+ This means $\vec{a} \cdot \vec{\nabla}\Phi = 0$ for any tangent vector to the surface \Rightarrow surface is equipotential

(has constant potential)

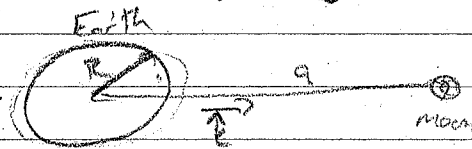


• Over long times, even the earth is effectively fluid, so the planet's surface is equipotential

● Tides:

- There is a daily variation in height of oceans due to moon's (and, not quite as much, the sun's) gravity

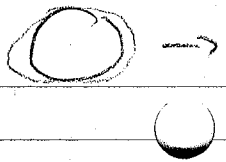
• In shorthand, the force on earth due to moon differs on opposite sides of the earth



• The "tidal force" is

$$\Delta\vec{F} = \vec{F}_{\text{moon}} - \vec{F}_c = \frac{GM_e M_m}{(a-R)^2} \hat{e} - \frac{GM_e M_m}{(a+R)^2} \hat{e} \approx \frac{4GM_e M_m R}{a^3} \hat{e}$$

- The near oceans are pulled toward the moon more than the rocky earth, which is pulled more toward moon than far oceans.
 \Rightarrow 2 high tides per day



- Potential Treatment.

- Treat the moon as a point object at position $\vec{a} \equiv a\hat{e}$ from the center of earth

+ The potential at any point \vec{r} on earth's surface is

$$\Phi(\vec{r}) = -\frac{G \cdot M_m}{|\vec{r} - \vec{a}|} = -\frac{G M_m}{[r^2 + a^2 - 2ar \cos \theta]^{1/2}}$$

+ Since $r \ll a$, we expand

$$\Phi \approx -\frac{G M_m}{a} \left[1 + \frac{r}{a} \cos \theta - \frac{1}{2} \left(\frac{r}{a}\right)^2 + \frac{3}{2} \left(\frac{r}{a} \cos \theta\right)^2 + \dots \right]$$

from $(1+x)^{-1/2} \approx 1 - \frac{1}{2}x + \frac{3}{8}x^2 + \dots$

+ keeps only linear term for $x = (r/a) \cos \theta$

+ First term is a constant. 2nd term

$$\Phi \approx -G M_m \frac{r}{a^2} \cos \theta \Rightarrow \text{uniform pull toward moon}$$

The last varies with position over the earth, leads to force partly inward along x, y plane (tidal force) + toward/away from moon

• Importance of the tidal forces:

+ First, think of a frame of earth accelerating toward moon to remove that force

+ Then, compared to gravitational field of earth,

$$\frac{\frac{\nabla \Phi}{M_m}}{\frac{GM_E}{R_E^2}} \approx \frac{M_m R_E / a^3}{M_E / R_E^2} = \frac{M_m R_E^3}{M_E a^3} \approx 5.7 \times 10^{-8}$$

- + Tidal force from sun is about $\frac{1}{2}$ as big \leftarrow coincidence
- + This is much smaller than various effects on earth, like centrifugal force. But they are noticeable b/c they change over the course of the day & month, i.e., 2 high tides per day of different sizes over the month.

• Estimated Heights of tides

- + Strict assumptions: earth is perfectly rigid sphere covered in water, which is always in equilibrium
- + Including the gravitational potential from the earth, the water experiences potential (ignoring centrifugal effects)

$$\Phi = \left[gh(\theta) - \frac{GM_{\oplus}}{R_{\oplus}} \right] - \frac{GM_m}{a^3} (R_{\oplus} + h(\theta))^2 \left(\frac{3}{2} \cos^2 \theta - \frac{1}{2} \right)$$

where $h(\theta) =$ height w.r.t. θ .

- + The surface of the water should be at constant potential.
For $h(\theta) \ll R_{\oplus}$

$$h(\theta) \approx \frac{GM_m R_{\oplus}^2}{g a^3} \left(\frac{3}{2} \cos^2 \theta - \frac{1}{2} \right) = \frac{M_m (R_{\oplus})^3}{M_{\oplus} (a)^3} \left(\frac{3}{2} \cos^2 \theta - \frac{1}{2} \right)$$

- + The scale is $\frac{M_m}{M_{\oplus}} \left(\frac{R_{\oplus}}{a} \right)^3 R_{\oplus} \approx 0.36 \text{ m}$ for the moon.

This seems small, but is reasonable for mid-ocean. Tides are greatly affected by the continental shelves and also possible resonances in places.

• Shapes of potentials for general matter distributions What is the potential like when mass is not spherically symmetric?

- The potential far away: answer usually complicated, but simplifies + has universal features at long distance

- The potential from a general distribution of mass at point \vec{r} is

$$\Phi(\vec{r}) = -G \int d^3r' \frac{\rho(\vec{r}')}{|\vec{r} - \vec{r}'|} = -G \int d^3r' \frac{\rho(\vec{r}')}{(r^2 + r'^2 - 2\vec{r} \cdot \vec{r}')^{1/2}}$$

+ Suppose $\rho(\vec{r}') = 0$ for $|\vec{r}'| < R$ (same fixed distance)
 then consider only points $|\vec{r}| \gg R$.

+ We can expand

$$\Phi(\vec{r}) = -G \int d^3r' \frac{\rho(\vec{r}')}{r} \left[1 - \frac{2\vec{r} \cdot \vec{r}'}{r^2} + \frac{r'^2}{r^2} \right]^{-1/2}$$

$$\equiv \Phi_0(\vec{r}) + \Phi_1(\vec{r}) + \Phi_2(\vec{r}) + \dots$$

• The contributions:

+ $\Phi_0(\vec{r}) = -GM/r$, where $M = \text{total mass}$

+ $\Phi_1(\vec{r}) = -2 \sum_i \vec{r} \cdot \left(\int d^3r' \rho(\vec{r}') \vec{r}' \right)_i \cdot \vec{r} = -\frac{GM}{r^3} \vec{r} \cdot \vec{R}$
 where $\vec{R} = \text{CM position}$. We can always make this term vanish (for gravity) by shifting origins

+ In EM, Φ_1 is the dipole term that can be nonvanishing if the total charge is zero.

$$+ \Phi_2(\vec{r}) = -\frac{G}{r^5} \int d^3r' \rho(\vec{r}') \left[\frac{3}{2} (\vec{r} \cdot \vec{r}')^2 - \frac{1}{2} (r')^2 r^2 \right]$$

$$\equiv -\frac{G}{r^5} \vec{r} \cdot (\underline{Q} \vec{r})$$

for quadrupole moment tensor

$$Q_{ij} = \int dm' (3r'_i r'_j - (r')^2 \delta_{ij})$$

(This is clearly related to the inertia tensor.)

+ If $\rho(\vec{r}')$ is axially symmetric around the z axis,

$$Q_{xy} = Q_{xz} = Q_{yz} = 0, \quad Q_{zz} = -2Q_{xx} = -2Q_{yy} = Q$$

where

$$Q = \int dm' (2z'^2 - x'^2 - y'^2)$$

so

$$\Phi_2(\vec{r}) = -\frac{GQ}{2r^3} (3\cos^2\theta - 1) \text{ for polar angle } \theta.$$

- Effects on + of Earth

- The earth is oblate (bulged out in middle) [vs prolate] due to centrifugal force.

+ Suppose Earth's equatorial radius is R_0 and polar radius is $R_0(1-\epsilon)$. By measurement $\epsilon \approx 1/300$. Let's understand why.

+ The centrifugal force of earth's rotation is $\frac{1}{2} m \omega^2 \rho^2$ on a mass m , so $\Phi_c = -\frac{1}{2} \omega^2 r^2 \sin^2 \theta$. cylindrical
The earth's surface, over geologic times, should become an equipotential, but $\Phi_0 + \Phi_c \neq \text{const}$. There must also be a quadrupole

+ Treat the oblate spheroidal shape of the Earth as an ellipse of revolution with $\epsilon \ll 1$. Then

$$\frac{r^2 \sin^2 \theta}{R_0^2} + \frac{r^2 \cos^2 \theta}{R_0^2 (1-\epsilon)^2} = 1 \Rightarrow r = R_0 (1 - \epsilon \cos^2 \theta)$$

Further the quadrupole is

$$Q = -\frac{4}{5} M R_0^2 \epsilon \quad \text{assuming uniform density}$$

+ Now we have $\Phi = \Phi_0 + \Phi_2 + \Phi_c \in \text{const}$, so the $\omega^2 \theta$ terms must cancel. We set

$$-\frac{GM}{R_0} \epsilon + \frac{3}{5} \frac{GM R_0^2}{R_0^3} \epsilon + \frac{1}{2} \omega^2 R_0^2 = 0$$
$$\Rightarrow \epsilon \approx \frac{5 \omega^2 R_0^3}{4g}$$

where $g = GM/R_0^2$ is gravitational acceleration

This number is a little too large b/c earth's core is more dense

• Some effects

+ The quadrupole force is not central, so the orbits of satellites precess around the earth's rotational axis.

+ The oblateness of the sun also causes the orbits of the planets to precess slightly. For Mercury, there is also a (measurable) but $\sim 10\times$ smaller effect from GR.